

## Estimating time lags in a simple coupled climate-carbon cycle model

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The carbon dioxide exchange between the atmosphere, ocean, and land, in presence of periodic external forcing of different type is considered. The aim is to determine the delay between global temperature, concentration of atmospheric, and oceanic carbon stock.

The exchange of CO<sub>2</sub> between the atmosphere and the ocean is described by a Bacastow-type model but with a temperature-dependent chemical constants in the ocean. CO<sub>2</sub> fluxes from atmosphere to land ecosystems and ocean are determined by

$$G = \beta_L(q - q_0) + \gamma_L T ,$$
$$F = F_0 \chi (q - q_0 - \zeta D q_0 / D_0) ,$$

where T is deviation of global temperature from the initial value, q is concentration of CO<sub>2</sub> in the atmosphere, D is deviation of carbon stock at the ocean from its initial value  $D_0 = 1.5 \cdot 10^5$  PgC,  $q_0 = 278$  ppm,  $\beta_L = (0.01 - 0.02)$  [GtC/ppm],  $\gamma_L = 0.05$  [GtC/Kyear], and  $F_0 = (2.5 - 4.5)$  [GtC/year] are coefficients,  $\chi$  is the solubility of CO<sub>2</sub> in the sea water,  $\zeta$  is evasion factor.

The coupled climate-carbon system is governed by equations:

$$C_0(dq/dt) = E(t) - F - G ,$$
$$dD/dt = F ,$$
$$C(dT/dt) = p \ln(q/q_0) + R(t) - \lambda T ,$$

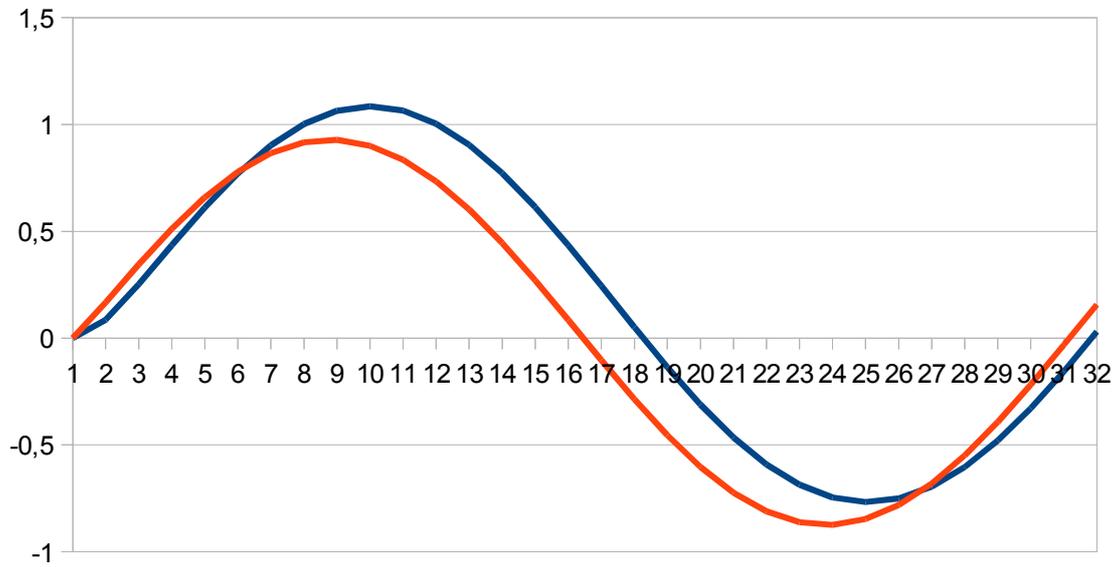
where E(t) is external (e.g., anthropogenic) CO<sub>2</sub> emissions to the atmosphere, R(t) is temperature forcing,  $p = 5.34$  [J/m<sup>2</sup> year],  $C_0 = 2.123$  [GtC/ppm], C is heat capacity per unit area ( $6 \cdot 10^7$  J/m<sup>2</sup> K),  $\lambda$  is feedback factor (1.37 W/m<sup>2</sup> K).

The temperature and emission forcing for simplicity are in the form of a sinusoid with the period from one thousand to 10 thousand years. Forcing of this type produces oscillations of T, q, and D. We determined mutual delays between these variables by maxima of the lagged correlation between them.

When temperature forcing is applied, q lags T by about one hundred years. If the period of the external forcing is changed, the lag increases approximately in proportion to this period. At an emission forcing T is late concerning q for 1 year (a time step of the model).

## References:

1. Boer G.J. and Arora V.K. (2013). Feedbacks in Emission-Driven and Concentration-Driven Global Carbon Budgets. *J. Climate*, **26**, 3326-3341.
2. Meier-Reimer E. and Hasselmann K. (1987). Transport and storage of CO<sub>2</sub> in the ocean – an inorganic ocean-circulation carbon cycle model. *Climate Dynamics*, **2**, 63-90.



**Figure 1.** Time series of T [K] (red line) and  $q/q_0 - 1$  (blue line) . Time is in centuries.